RADIOCARBON SIMULATION FAILS TO SUPPORT THE SYNCHRONEITY REQUIREMENT OF THE YOUNGER DRYAS IMPACT HYPOTHESIS

Supplementary Data: Extended Methods and Alternative Simulations

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1. Detailed Description of Simulation Steps

The main text provides a general outline of the simulation, while this section describes the simulation steps in detail. A more detailed overview of the simulation can be obtained by downloading the R script and the comments contained within. Section 2 provides instructions for running this script.

The simulation is performed for a unique event, *E*. Here, *E* corresponds to either the hypothesized Younger Dryas Impact or the Laacher See volcanic eruption. For each event *E*, we performed the simulation for 10,000 iterations in each year $x_{E,i}$ in a vector x_E of calendar years. For the former event, we used dates corresponding to the purported Younger Dryas Boundary (YDB), and for the latter event we used dates associated with the Laacher See Tephra (LST). For the YDB, x_{YDB} spans 151 years within 12,860–12,710 cal yr BP, and for the LST, x_{LST} spans 151 years within 12,991–12,841 cal yr BP. The central 101 years in each x_E vector corresponds to a range of possible calendar years for each hypothesized event. We simulated over a 25-year buffer on each side of the hypothesized range to observe how expectations vary for calendar years proximate to the possible calendar years for each hypothesized event.

For each of the 10,000 iterations in each year $x_{E,i}$, the simulation performs four main steps:

- (1) Create a vector of true calendar ages *C* for the synchronous event. The length of *C* depends on the number of ¹⁴C measurements in the observed sample (30 for the YDB and 19 for the LST).
- (2) Generate expected *target* ¹⁴C values for each true calendar age C_i in vector C.
- (3) For each *target* ¹⁴C value, generate a *measured* ¹⁴C value as observed by a laboratory.
- (4) Calibrate each *measured* ¹⁴C value with the IntCal13 ¹⁴C calibration curve (Reimer et al., 2013).

The remainder of this section describes each of the four steps. These descriptions concern a single iteration for a single calendar age $x_{E,i}$. To help illustrate these steps, Table S1.1 provides a toy set of observed ¹⁴C measurements that could be used as simulation input for a hypothetical synchronous event. We expand on this table with each simulation step. All references to "calendar ages/years/dates" indicate years BP (AD 1950). AMS refers to accelerator mass spectrometry, GPC refers to gas proportional counting, and LSC refers to liquid scintillation counting.

1.1. True Calendar Ages

For each iteration within calendar year $x_{E,i}$, a vector of repeated $x_{E,i}$ values is first generated. The length of this vector is determined by the number of measurements in the observed ¹⁴C dataset. The YDB dataset consists of 30 reported dates, which creates a vector of 30 $x_{E,i}$ values for these simulations. The corresponding LST vector length is 19. For simulations that lack "old wood" effects, these repeated vectors are treated as the true calendar ages for each simulated ¹⁴C measurement.

Some versions of the simulation account for "old wood" effects with an "old wood" model (OWM). These versions add a calendar year offset to each $x_{E,i}$ value. Additional years are added only to the $x_{E,i}$ values that correspond to dated materials from organisms that may have died prior to the event of interest. Twenty-four of 30 reported YDB measurements originate from such materials, and 18 of 19 reported LST measurements correspond to these materials. Short-lived samples, such as those from grasses or seeds, do not receive age offsets. The new vector of calendar years that includes offset values is referred to here as x_E^o (Table S1.2).

Table S1.1. A toy dataset of reported ¹⁴C measurements for five samples associated with a hypothesized event E. These measurements were made by three laboratories with codes ABCD, EFGH, and IJKLM. The first and last laboratories measure ¹⁴C via AMS, while EFGH measures ¹⁴C via GPC. Three of the five measurements are on wood, indicating that they may correspond to calendar ages older than the event of interest. The simulation generates expected ¹⁴C values (i.e., what might be observed in the μ column given a true synchronous event), while all other columns in this table provide context dependent input for the simulation itself.

	Reporte	d ¹⁴ C		
Sample ID	μ	σ	"old wood"?	Lab type
ABCD-0001	10,251	25	No (seed)	AMS
ABCD-0002	10,290	35	No (seed)	AMS
EFGH-0001	10,295	75	Yes (wood)	GPC
EFGH-0003	10,299	55	Yes (wood)	GPC
IJKLM-0001	10,321	40	Yes (wood)	AMS

OWM age offsets are drawn randomly from an exponential distribution, with offset values near zero more likely. In other words, OWM offsets assume that organism death most likely occurred shortly before the calendar year of the event *E*, with the probability of an earlier death event decreasing with temporal distance from event *E*. We considered two exponential distributions, one with the λ parameter set to 0.04 and one with λ set to 0.01 (Figure S1.1). The former distribution is a conservative scenario in which old samples predate the event by few years ($\mu = 25$, 95% HDI = 0–75), while the latter distribution results in larger age offsets on average ($\mu = 100$, 95% HDI = 0–300; Figure S1.1b,c). New values are drawn from each exponential distribution for each of the 10,000 iterations, allowing for λ specific variability in "old wood" effects to be estimated across 10,000 iterations. These distributions were selected to bound the extremes of realistic offsets that might be expected given "old wood" effects.

Table S1.2. The initial step in a single simulation iteration. Here, the toy dataset is used as input for a simulation iteration at 12,000 cal yr BP. The reported ¹⁴C means have been removed, as these do not serve as input for the simulation. The age offsets are shown for hypothetical values drawn from an exponential distribution with $\lambda = 0.04$.

	Report	ted ¹⁴ C				OWM	
Sample ID	μ	σ	"old wood"?	Lab type	X E,i	Offset	Х Е,i ^O
ABCD-0001	-	25	No	AMS	12,000	N/A	12,000.00
ABCD-0002	-	35	No	AMS	12,000	N/A	12,000.00
EFGH-0001	-	75	Yes	GPC	12,000	27.5	12,027.50
EFGH-0003	-	55	Yes	GPC	12,000	6.10	12,006.10
IJKLM-0001	-	40	Yes	AMS	12,000	88.50	12,088.50

Most of the wood samples for Laacher See are *Populus* (Baales et al., 2002), which is most likely Eurasian aspen (*Populus tremula*) or black poplar (*Populus nigra*). The former species lives an average of 50–100 years (Caudullo and de Rigo, 2016) and can live up to 200 years (von Wühlisch, 2009). In contrast, black popular generally live only 20–50 years. As such, the conservative exponential model ($\lambda = 0.04$) generally gives offsets intermediate between the lifespans of these species (Figure S1.1b). The latter "old wood" exponential function ($\lambda = 0.01$) is more consistent with aspen and produces occasional offsets several centuries older than the lifespan of aspen (Figure S1.1c). This can accommodate scenarios in which dead wood on the

landscape is incorporated into the geological stratum of interest. We assume that true "old wood" effects for the YDB and LST wood and charcoal ¹⁴C samples fall somewhere between the two extremes defined by these exponential functions, and neither exponential function is intended to match precisely the "old wood" effects in either context, which are probably unknowable.



Figure S1.1. Exponential distributions used for versions of the simulation employing an OWM. (a) Probability density functions for each offset. (b and c) 30 examples of 20 "old wood" dates for a synchronous event at 12,000 cal yr BP, where each row is a sample of 20 dates. Panel b corresponds to an exponential distribution with $\lambda = 0.04$, and panel c corresponds to an exponential distribution with $\lambda = 0.04$.

1.2. Target ¹⁴C Values

We generated a target ¹⁴C value for each $x_{E,i}^{O}$ value in two steps. First, mean ¹⁴C values were obtained by "uncalibrating" $x_{E,i}^{O}$ values with the uncalibrate function in the *rcarbon* R package (Bevan and Crema, 2018). This function works by drawing from a normal distribution of ¹⁴C values associated with a calendar year, where the distribution corresponds to the calibration curve error in the IntCal13 ¹⁴C calibration curve (Figure S1.2a). This produces a hypothetical mean atmospheric ¹⁴C value since they correspond to a calendar year with the same uncalibrated mean atmospheric ¹⁴C value since they correspond to a calendar year with the sample hypothetical mean atmospheric ¹⁴C value. Mean atmospheric ¹⁴C values are resampled for each of the 10,000 iterations.

Second, we account for intra-annual variability around the mean atmospheric ¹⁴C value in a calendar year. We estimated the difference between seasonal extremes of atmospheric ¹⁴C variability using data from McDonald et al. (2019). They estimate the distance between seasonal extremes of atmospheric ¹⁴C with two calculation methods, each of which is performed for situations in which atmospheric ¹⁴C production is in increasing and decreasing states (Table S1.4). We also consider stable atmospheric ¹⁴C production, treated here as the midpoint between McDonald et al.'s (2019) values for increasing and decreasing production.

Table S1.3. The first part of step two for a single simulation iteration. Here, the toy dataset is used as input for a simulation iteration at 12,000 cal yr BP. The reported ¹⁴C means have been removed, as these do not serve as simulation input. Mean atmospheric ¹⁴C values have been sampled from the IntCal13 ¹⁴C calibration curve error distribution around each calendar age $x_{E,i}^{O}$.

	Report					
Sample ID	μ	σ	"old wood"?	Lab type	X E,i ^O	Mean atmospheric ¹⁴ C
ABCD-0001	-	25	No	AMS	12,000.00	10,222
ABCD-0002	-	35	No	AMS	12,000.00	10,222
EFGH-0001	-	75	Yes	GPC	12,027.50	10,261
EFGH-0003	-	55	Yes	GPC	12,006.10	10,300
IJKLM-0001	-	40	Yes	AMS	12,088.50	10,366



Figure S1.2. A visual schematic of simulated target ¹⁴C values based on context dependent input from the toy dataset: (a) "Uncalibrating" five calendar ages (i.e., converting x_{E,i}^O values into mean atmospheric ¹⁴C values; Table S1.3 for details). The blue calibration curve bands show the 50% and 95% error regions, and the white line shows the mean value of the curve. (b) 30 randomly sampled beta distributions that represent possible intra-annual distributions of ¹⁴C variability. (c) Mean "uncalibrated" ¹⁴C values (dotted lines), intra-annual ¹⁴C variability around those values based on 10,000 randomly sampled beta distributions (red semitransparent regions), and ¹⁴C values sampled from those intra-annual ¹⁴C variability distributions (solid lines). The ¹⁴C values sampled from each intra-annual distribution comprise the target ¹⁴C values. These are the values that laboratories attempt to measure. In panels a and c, light red geometry corresponds to three calendar ages with "old wood" effects, and dark red geometry corresponds to two overlapping calendar ages that date the event of interest.

Table S1.4. Estimates of the intra-annual distance between atmospheric ¹⁴C extremes (McDonald et al., 2019 for increasing and decreasing values. The stable values are the midpoints between the increasing and decreasing values).

Intra-annual distance	Error	Atmospheric ¹⁴ C trend	Calculation method
18.0	13.0	Increasing	First
23.0	16.0	Decreasing	First
20.5	14.5	Stable	First
26.0	16.0	Increasing	Second
22.0	13.0	Decreasing	Second
24.0	14.5	Stable	Second

To estimate intra-annual atmospheric ¹⁴C variability, we first modelled possible values as sparsely beta distributed $\beta(2, 2)$, centered on 0, and scaled by the intra-annual distance. Since the intra-annual distance is imprecisely known, we simulated 10,000 beta distributions scaled by samples drawn from N(Intra-annual distance, Error) (Figure S1.2b). For each of the 10,000 samples, we randomly selected one of the six possible distance and error pairings. Therefore, these distributions average across atmospheric ¹⁴C production trends and calculation methods. For each of the 10,000 simulation iterations, a random value is drawn from one of the 10,000 centered and scaled beta distributions. This random value is then added to the mean atmospheric ¹⁴C value to estimate a target ¹⁴C value that reflects both inter- and intra-annual ¹⁴C variability (Table S1.5; Figure S1.2c).

Table S1.5. The second part of step two for a single simulation iteration. Here, the toy dataset is used as input for a simulation iteration at 12,000 cal yr BP. The reported ¹⁴C means have been removed, as these do not serve as input for the simulation.

	Repo	Reported ¹⁴ C "old		Reported ¹⁴ C "old Lab		Mean	Mean Intra-annual		
Sample ID	μ	σ	wood"?	type	atmospheric ¹⁴ C	variation offset	Target ¹⁴ C		
ABCD-0001	-	25	No	AMS	10,222	15	10,237		
ABCD-0002	-	35	No	AMS	10,222	-1	10,221		
EFGH-0001	-	75	Yes	GPC	10,261	-6	10,255		
EFGH-0003	-	55	Yes	GPC	10,300	-1	10,299		
IJKLM-0001	-	40	Yes	AMS	10,366	1	10,367		

1.3. Measured ¹⁴C Values

Given minor inter-laboratory variability that conditions systematic biases, as well as minor instrumental error that shapes intra-laboratory measurement repeatability, measured ¹⁴C values depart from their target values (Boaretto et al., 2003; International Study Group, 1982; Scott et al., 1990, 1998, 2010a; Scott, Cook, and Naysmith, 2007). As such, the dispersion of reported ¹⁴C values should vary based on the number of laboratories that have contributed to a reported ¹⁴C dataset, as well as based on laboratory specific characteristics that might further influence the repeatability of measurements. These sources of inter- and intra-laboratory variability can be used to model expected departures of ¹⁴C measurements from their target values, given known numbers of laboratories and measurements per laboratory.

To estimate this variability, we fit a Bayesian multilevel model to data reported in the Fifth International Radiocarbon Intercomparison (VIRI) (Scott et al., 2010a, 2010b; Scott, Cook, Naysmith et al., 2007). We refer to this model as the Laboratory Measurement Bias and Repeatability Model, or LBM. The LBM treats ¹⁴C measurements as a normally distributed outcome. We defined a three-parameter linear model for the outcome mean: a categorical intercept for each sample material, a random categorical effect for laboratory ID, and a scaling parameter that adjusts the random laboratory ID effect based on whether the laboratory performs AMS or GPC/LSC measurements. We defined a four-parameter linear model for the outcome standard deviation. Parameters consist of a baseline intercept, a random categorical effect for laboratory ID, a linear effect for reported measurement error, and a categorical effect that accounts for whether laboratories perform AMS or GPC/LSC measurements. Section 3 describes this model in detail, including the model formula, prior distributions for model parameters, and a posterior predictive check.

The simulation samples 10,000 sets of values from the posterior distributions of the LBM parameters, with a unique set of parameter values applied to each of the 10,000 simulation iterations (Figure S1.3). Therefore, uncertainty in the model parameters is distributed across iterations within each calendar year $x_{E,i}$.

- 1. A random offset from the target ¹⁴C value representing the mean observed value within each laboratory. The number of sampled offsets is determined by the number of laboratories that contributed to the reported ¹⁴C dataset (Table S1.6).
- 2. A multiplier term that rescales the offset for GPC/LSC laboratories (Table S1.6).
- 3. A random within-laboratory standard deviation that is rescaled by an additional multiplier value for GPC/LSC laboratories (σ_L). This standard deviation further varies by the error reported for each ¹⁴C measurement (Table S1.7).
- Values are then drawn from laboratory and sample specific distributions defined by N(Mean laboratory specific ¹⁴C, σ_L), representing ¹⁴C values that might be measured by each laboratory (Table S1.8).

The standard deviation for expected within lab measured ¹⁴C variability depends on the reported error of the sample (σ). It takes the form,

$$\sigma_L = exp(Lab \ effect + GPCLSC \ effect + log \ (\sigma) * \sigma \ effect) * 100.01.$$

(Equation S1.1)

Possible values for each effect are presented in Table S1.7. The 100.01 value to the right of the exponential transformation puts the result on the scale of ¹⁴C years (the LBM is fitted to ¹⁴C year z-scores, and therefore, the output of this linear model needs to be put back on the ¹⁴C year scale).

The measured ¹⁴C values comprise an expected set of observations generated for a single iteration. The simulation records the standard deviation of these values, σ^{14} C (69.80 for the toy dataset detailed here), as an expected measure of dispersion for a series of ¹⁴C measurements, given a synchronous event. This is completed over 10,000 iterations for each calendar year $x_{E,i}$, yielding a distribution of simulated σ^{14} C values given the number of labs, lab types (AMS or GPC/LSC), reported measurement errors, and potential "old wood" effects associated with a reported ¹⁴C dataset. For versions of the simulation that exclude the LBM, these σ^{14} C values are calculated with the target ¹⁴C values rather than the measured ¹⁴C values (as calculated in Step 2).



Figure S1.3. An example of five target ¹⁴C values being converted to measured ¹⁴C values via sampling from the LBM. Refer to Tables S1.6, S1.7, and S1.8 for the values depicted in this figure. Light red lines correspond to three calendar ages with "old wood" effects, and dark red lines correspond to two calendar ages that date the event of interest. Rotated normal distributions illustrate intra-laboratory sampling distributions for each measured ¹⁴C value; They correspond to the distribution of repeated measurements for a given laboratory at a given reported measurement error, centered on the mean laboratory specific ¹⁴C value for a target ¹⁴C value. Two normal distributions with empty fills and dark red outlines depict calendar ages that lack "old wood" effects, and three normal distributions with light red fills and no outlines show calendar ages with "old wood" effects. Note, the intra-laboratory sampling distributions for ABCD-0001 and EFGH-0001 overlap nearly completely, and the values sampled for ABCD-002 and EFGH-0001 are very similar.

Table S1.6. Obtaining mean laboratory offset values for measured ¹⁴C in a single simulation iteration. Here, the toy dataset is used as input for a simulation iteration at 12,000 cal yr BP. The reported ¹⁴C means have been removed, as these do not serve as input for the simulation. The laboratory offset and GPC/LSC multiplier values are sampled from the LBM parameters, which vary across the 10,000 simulation iterations.

	Repor	ted ¹⁴ C	_	Target	Lab	GPC LSC	Mean lab
Sample ID	μ	σ	Lab type	¹⁴ C	offset	multiplier	specific ¹⁴ C
ABCD-0001	-	25	AMS	10,237	12.50	N/A	10,249.50
ABCD-0002	-	35	AMS	10,221	12.50	N/A	10,233.50
EFGH-0001	-	75	GPC	10,255	-5.80	1.1	10,248.60
EFGH-0003	-	55	GPC	10,299	-5.80	1.1	10,292.60
IJKLM-0001	-	40	AMS	10,367	22.00	N/A	10,389.00

Table S1.7. Parameters for sampling measured ¹⁴C values. Here, the toy dataset is used as input for a simulation iteration at 12,000 cal yr BP. The reported ¹⁴C means have been removed, as these do not serve as input for the simulation. Laboratory effects, GPC/LSC effects, and σ effects are sampled from the posterior distributions of the LBM parameters, which vary across the 10,000 simulation iterations.

Reported ¹⁴ C		Lab Mean lab		Within lab std. deviation (σ _L) parameters			
Sample ID	μ	σ	type	specific ¹⁴ C	Lab effect	GPC LSC effect	σ effect
ABCD-0001	-	25	AMS	10,249.50	-2.10	0	0.35
ABCD-0002	-	35	AMS	10,233.50	-2.10	0	0.35
EFGH-0001	-	75	GPC	10,248.60	-2.50	0.32	0.35
EFGH-0003	-	55	GPC	10,292.60	-2.50	0.32	0.35
IJKLM-0001	-	40	AMS	10,389.00	-1.90	0	0.35

Table S1.8. Sampled measured ¹⁴C values. Here, the toy dataset is used as input for a simulation iteration at 12,000 cal yr BP. The reported means have been removed, as these do not serve as input for the simulation. Lab/sample deviations were calculated with the "Within lab std. deviation parameters" from Table S1.7 using Equation S1.1.

	Reported ¹⁴ C μ σ		Lab	Within lab s	Measured ¹⁴ C	
Sample ID			type	¹⁴ C μ	Std. deviation (σ _L)	value
ABCD-0001	-	25	AMS	10,249.50	18.40	10,241.20
ABCD-0002	-	35	AMS	10,233.50	19.10	10,251.70
EFGH-0001	-	75	GPC	10,248.60	18.90	10,250.80
EFGH-0003	-	55	GPC	10,292.60	18.40	10,296.30
IJKLM-0001	-	40	AMS	10,389.00	23.60	10,408.60

1.4. Calibrated ¹⁴C Measurements

The simulation then calibrates the measured ¹⁴C values using the errors described in the reported dataset, producing a probability density across calendar ages for each measurement (Table S1.9; Figure S1.4). This is accomplished with the IntCal13 curve using the calibrate function in the *rcarbon* R package (Bevan and Crema, 2018).

Table S1.9. Sampled measured ¹⁴C values and the 95% highest density intervals (HDI) for their calibrated age densities. Here, the toy dataset is used as input for a simulation iteration at 12,000 cal yr BP. The reported means have been removed, as these do not serve as input for the simulation. Note that the measured ¹⁴C values in the simulation are calibrated with σ values for the reported measurements.

Reported ¹⁴ C		Measured ¹⁴ C				
Sample ID	μ	σ	μ	σ	Cal yr BP (95% HDIs)	
ABCD-0001	-	25	10,241.20	25	12,111–11,921; 11,915–11,827	
ABCD-0002	-	35	10,251.70	35	12,130–11,826	
EFGH-0001	-	75	10,250.80	75	12,384–12,264; 12,246–11,711	
EFGH-0003	-	55	10,296.30	55	12,386–12,262; 12,249–11,929; 11,894–11,829	
IJKLM-0001	-	40	10,408.60	40	12,517–12,482; 12,424–12,085	



Figure S1.4. Toy dataset: Five simulated ¹⁴C measurements (rotated normal distributions) calibrated with the IntCal13 calibration curve (white line with blue bands). The blue calibration curve bands show the 50% and 95% error regions and the white line depicts the mean curve value. Light red geometry corresponds to three samples with "old wood" effects, and dark red geometry corresponds to two overlapping samples that date the event of interest. The undulating distributions on the x-axis depict calibrated age densities, which are used to calculate dissimilarity values.

Following calibration, dissimilarity values are then obtained by first calculating the Manhattan distance between each pair of age densities (Table S1.10):

Manhattan distance =
$$\sum_{i=1}^{c} |A_{x,i} - A_{y,i}|$$
,

(Equation S1.2)

where $A_{x,i}$ indexes the proportion of age density A_x in calendar year *i*, $A_{y,i}$ indexes the proportion of age density A_y in calendar year *i*, and *c* is the length of a vector defined by the union of calendar ages shared by the pair of age densities. The expected dissimilarity value is then calculated by taking the mean of all pairwise Manhattan distances and dividing this mean by two. A value of exactly zero indicates that the age densities are identical, while a value of exactly one indicates that the set of age densities are completely nonoverlapping. Like the σ^{14} C values calculated at the end of Step 3, 10,000 dissimilarity values are obtained for each year $x_{E,i}$ across the iterations.

Table S1.10. Manhattan distances between each pair of the five calibrated age distributions in the
toy dataset. The mean of these values is divided by two to obtain a measure of dissimilarity (0.508
for this toy dataset). Here, the toy dataset is used as input for a simulation iteration at 12,000 cal
yr BP.

<u>y' 0' '</u>					
	ABCD-001	ABCD-002	EFGH-001	EFGH-003	IJKLM-001
ABCD-001					
ABCD-002	0.19				
EFGH-001	0.71	0.62			
EFGH-003	0.95	0.80	0.63		
IJKLM-001	1.86	1.81	1.48	1.14	

2. Guide to Running the Simulation

This simulation consists of an R script that is designed to run on a cluster. Parallelization occurs across the range of 302 simulated years (151 years each for the LST and the YDB). Optimally, one year is assigned per available core, plus one additional core for the master process. If at least one core is available per year, execution time is mainly limited by the number of simulation iterations. The results in this paper were obtained by executing the R script on the ManeFrame II cluster at Southern Methodist University (SMU), utilizing one core per simulated year. Under these circumstances, a 10,000-iteration simulation completes in about 36 hours, a 1000-iteration simulation sunder 1000 iterations generally provide noisy output distributions of σ^{14} C and dissimilarity values. If cores are limited and multiple years are run per core, execution time will increase dramatically depending on the core that is assigned the most years over which to simulate (henceforth, maximum-core-years). Estimated execution time is roughly maximum-core-years is three for a 10,000-iteration simulation on the ManeFrame II cluster, the expected time is 3*36 hours.

2.1. Requirements

<u>Stan</u>: The LBM is fit via Hamiltonian Monte Carlo simulation in Stan (Stan Development Team, 2018). Visit <u>https://www.mc-stan.org</u> for installation details.

<u>R packages</u>: *rstan* (Stan Development Team, 2018), *ggplot2* (Wickham, 2016), *parallel* (R Core Team, 2018), *reshape2* (Wickham, 2007), *rcarbon* (Bevan and Crema, 2018), *matrixStats* (Bengtsson et al., 2018), *patchwork* (Pedersen, 2018), and *rethinking* (McElreath, 2017). The first six packages are available in the CRAN.

patchwork is available at: <u>https://github.com/thomasp85/patchwork</u> rethinking is available at: <u>https://github.com/rmcelreath/rethinking</u>

<u>Data files</u>: *IRI.csv* (table of reported ¹⁴C measurements from the Fifth International Radiocarbon Intercomparison), *RCmeasurements.csv* (table of dates reported for the Laacher See Tephra and Younger Dryas Boundary).

The script must be executed on a cluster with cores distributed across nodes. The main file output (*SimDat#.RData*) for the 10,000-iteration simulation is 587 MB, and there are minor memory spikes during the simulation (intermediate R objects are created during the simulation that are not included in the main output file).

2.2. Running the R Script

Place the R script, *IRI.csv*, and *RCmeasurements.csv* in your working directory. Ensure that Stan and all required R packages are installed. Follow these steps:

- 1. <u>Adjust user-arguments for simulation</u>. Navigate to the 'USER ARGUMENTS' block of code in the R script to adjust the simulation as appropriate (code lines ~220-305). This block contains variables that specify the number of nodes to be used, the number of cores per node, the number of simulation iterations, plotting options, the ranges of calendar years over which to simulate each event, OWM parameter values, and other variables of interest. Inline code comments further detail each variable.
- 2. <u>Source the R script</u>. This will load the csv data files and prepare parameters for the simulation. On the first run, the script fits the LBM, which may take 5-15 minutes. After the

model is fitted, posterior parameter values are exported in RData file *IRI#.RData* (where '#' is the number of user-specified iterations for the simulation). If you leave this file in your working directory and you plan to run the simulation again in the future, the script will read *IRI#.RData* into the simulation rather than refit the model, saving run time. Every time a simulation is run with a new number of iterations (i.e., new '#' values), the model will be refitted. As such, users can store multiple *IRI#.RData* files for running the simulation with different numbers of iterations. The correct file will be read automatically for each simulation if it has already been generated in the working directory.

- 3. <u>Monitor the working directory for intermediate output files</u>. In addition to *IRI#.RData*, the script will output *SimDatIntermediate.RData*. This file contains simulation parameters to be read by each node. If you wish to delete this file, do not so until the simulation has initiated on every node (i.e., every node has imported the simulation parameters from *SimDatIntermediate.RData*). Immediately following the creation of *SimDatIntermediate*.*RData*, the main R script will create daughter R script files with the filename *NodeSim#.R*. The number of these scripts that is created corresponds to the number of user-specified nodes. After these scripts appear in the working directory, move to Step 4.
- 4. <u>Submit daughter scripts to the cluster</u>. The main R script also outputs an sbatch array submit script that can be executed to request nodes for every daughter R script (*nodesim.sh*). This is formatted to run on the ManeFrame II cluster at SMU, but it can be easily edited to run on other clusters using a Slurm workload manager. Alternatively, you may submit the daughter scripts to nodes using your own method.
- 5. <u>Wait for results</u>. After you submit the daughter scripts, the main R script waits for them to complete (it scans the working directory every 30 seconds for output from the daughter scripts). When a daughter script completes, it outputs *NodeDat#.RData*. After all daughter output files are present in the working directory, the main script automatically imports them, aggregates the results, and creates one output RData file (*SimDat#.RData*, where '#' is the number of user specified simulation iterations).

2.3. Results and Output Files

Simulation results are contained in objects stored in *SimDat#.RData*. Comments in the script describe these objects. This file is automatically written to the working directory after the simulation completes. *SimDat#.RData* can be opened and explored in an interactive R session on a personal computer. Although the intermediate output files may be of interest (*IRI#.RData*, *SimDatIntermediate.RData*, *NodeDat#.RData*, *NodeSim#.R*, and *nodesim.sh*), they do not contain the primary results and may be deleted after the simulation is completed.

3. Laboratory Measurement Bias and Repeatability Model (LBM)

This section has four parts: (1) a description of the Fifth International Radiocarbon Intercomparison (VIRI) dataset to which the LBM was fitted, (2) model formulas and prior distributions for parameters, (3) a description of the posterior parameter values, and (4) a description of Hamiltonian Monte Carlo (HMC) diagnostics with a posterior predictive check. The goal of this model is to estimate parameters that describe inter-lab variation in the measurement of the ¹⁴C content of a sample as well as intra-laboratory variation over repeated measurements of a sample. This involves estimating the distributions of mean ¹⁴C values and ¹⁴C standard deviations across laboratories.

3.1. The VIRI Dataset

To fit the model, we first aggregated data presented in the VIRI (Scott et al., 2010a, 2010b; Scott, Cook, Naysmith et al. 2007). These data included all ¹⁴C measurements across all sample materials for which ¹⁴C measurements were reported (samples B, D, F, G, H, and I). In total, this spans 420 measurements performed by 80 laboratories.

We transformed these data in three steps. First, we centered all ¹⁴C measurements on the median value for each sample material. Second, we used these centered measurements to investigate the presence of outliers. Outliers were conservatively defined as measurements that fall at least six times the interquartile range (IQR) distance outside the first and third quartile within each set of ¹⁴C measurements for a sample material. The values identified as outliers may have resulted from unusually poor quality-control for some laboratories or from other anomalies in measurement. After outliers were identified, all measurements associated with the laboratory that produced an outlier were removed from the dataset (Figure S3.1). This reduced the number of laboratories from 80 to 68, and the number of ¹⁴C measurements from 420 to 361. Finally, these median-centered measurements were converted to z-scores (Figure S3.1).

3.2. Model Formula and Prior Distributions for Parameters

Each of the 361 median-centered ¹⁴C z-scores is associated with four additional variables: a categorical sample material ID (B, D, F, G, H, or I), a categorical laboratory ID, a dummy variable indicating whether than laboratory performed an AMS (0) or GPC/LSC measurement (1), and the reported measurement errors for ¹⁴C values. Although measurement errors should represent uncertainty in the reported means, laboratories calculate these errors in a variety of ways that may not be comparable (Scott, Cook, and Naysmith, 2007). As such, we treat them as predictor variables for the dispersion of reported ¹⁴C means, with the expectation that within-laboratory ¹⁴C measurement dispersion increases with larger reported errors.

First, we defined the likelihood for median-centered ¹⁴C z-scores as

¹⁴ $C \sim N(\mu, \sigma)$.

(Equation S3.1)

Median-centered ¹⁴C z-scores are distributed N(μ , σ). We then modelled μ as a linear outcome of the ¹⁴C value of sample material i, C_s[sample i], and an offset from that sample material value that depends on the laboratory ID, C_o[lab j]. The laboratory ID offset also varies based on the AMS vs GPC/LSC dummy variable, AMS, through parameter C_{AMS}:

$$\mu = C_s[sample i] + C_o[lab j] \times (1 + (C_{AMS} - 1) \times AMS).$$
(E)

(Equation S3.2)

Since median-centered ¹⁴C z-scores likely approximate a true sample-specific value near zero, we use a prior distribution of N(0, 1) for each C_s[sample i]. We modelled C_o[lab j] as distributed N(0, C₀^{σ}), with the prior for C₀^{σ} set to exp(2). C_{AMS} may reduce or increase the effect of C_o[lab j], but it does not affect the sign of C_o[lab j], taking only positive values. Values less than 1 reduce the lab specific offset C_o[lab j], while values greater than 1 increase the lab specific offset C_o[lab j]. As such, for C_{AMS} we use a gamma distribution with the mean centered on 1 as a prior: gamma(1, 0.5). The parameterization of this prior does not follow the base *dgamma* R function, but instead uses the *dgamma2* parameterization included in the *rethinking* R package (McElreath, 2017).



Figure S3.1. Histograms of median-centered ¹⁴C z-scores for samples B, D, F, G, H, and I. Left panels show all available measurements (n = 420), and right panels show only those measurements that originate from laboratories that did not produce outlier values (n = 361). The values in the right panels were used to fit the LBM.

 σ can then be interpreted as intra-laboratory variation in median-centered ¹⁴C z-scores. We modelled σ as the linear outcome of a baseline parameter shared by all laboratories, σ_{I} , a

laboratory ID specific offset parameter, σ_{lab} [lab j], the log reported measurement error (log(ME)) multiplied by parameter σ_{ME} , and a parameter that is expressed only for GPC/LSC laboratories (σ_{AMS}). As such, the dispersion of repeated intra-laboratory ¹⁴C z-scores depends on laboratory specific variation in repeatability, the reported measurement error associated with those measurements, and whether the measurement was obtained via AMS or GPC/LSC. Laboratory specific variation depends on each reported ¹⁴C measurement error, which can vary from value to value within a single laboratory. To constrain σ on the positive scale, we used a log link function:

$$log(\sigma) = \sigma_{lab}[lab j] + \sigma_l + \sigma_{ME} \times log(ME) + \sigma_{AMS} \times AMS.$$

(Equation S3.3)

We assigned the same informative normal prior distribution to σ_{I} , σ_{I} , and σ_{AMS} : N(0, 1). We modelled $\sigma_{Iab}[Iab j]$ as distributed N(0, $\sigma_{Iab}{}^{\sigma}$), with the prior for $\sigma_{Iab}{}^{\sigma}$ set to exp(2). Readers may note that we have not modelled covariance between the random laboratory parameters, C_o[Iab j] and $\sigma_{Iab}[Iab j]$. For the simulation, the practical implication of this decision is that these parameter values are sampled independently for simulated laboratories rather than from a multivariate distribution. Such covariance is often modelled as multivariate normal, which would be inappropriate here, as σ_{Iab} should vary with only the magnitude of C_O rather than the magnitude and sign of C_O (i.e., a parabolic rather than monotonic relationship). In other words, we might expect within-laboratory dispersion to vary with absolute laboratory ID offset, regardless of whether the laboratory ID offset is above or below the target ¹⁴C value (Figure S3.2). For the sake of model simplicity and interpretability, we did not attempt to model a parabolic relationship between these parameters.



Figure S3.2. (a) Example relationships that might be expected if laboratory parameters covaried monotonically. (b) Example relationships that might be expected given a parabolic relationship between parameters. Given the lack of a covariance component in the model, laboratory parameters were estimated independently, conforming to the horizontal grey relationship in each panel (i.e., no covariance).

This model assumes that systematic offsets between mean laboratory measurements and target ¹⁴C values are maintained across any sample materials that a laboratory might measure. For example, consider a laboratory faced with measuring three different sample materials: A, B,

and C. If the laboratory takes measurements that are on-average -10 ¹⁴C years from the target value of sample A, this mean offset will be also be present when the same laboratory measures the ¹⁴C values of sample materials B and C. If the systematic offset varies between sample materials A, B, and C, this cannot be captured by the model. This model also assumes that within-laboratory repeatability is uniform across sample materials A, B, and C. In reality, different sample material types (e.g., bone, wood, grass seeds) and variability in target ¹⁴C values may affect both systematic offsets and within-laboratory repeatability. Unfortunately, the available VIRI dataset is insufficient to explore these issues in detail with this model.

The LBM was specified using the *map2stan* function in the *rethinking* R package (McElreath, 2017) and fitted through Hamiltonian Monte Carlo (HMC) simulation in *rstan* (Stan Development Team, 2018). The model was fitted with four chains, each of which performed 5000 warmup and 2500 sampling iterations (10,000 total sampling iterations).

3.3. Posterior Parameter Values

Posterior estimates for each sample material ¹⁴C value are close to each material's observed sample median (Figure S3.3a). Each posterior distribution for these sample materials includes its respective observed median value in a high-density region (since sample values are median centered, the observed median value is 0 for each distribution). The posterior distribution for C_{AMS} has a mean of 1.26, indicating that GPC/LSC laboratories generally have larger mean offsets from target ¹⁴C values than do AMS laboratories (Figure S3.3b). However, this posterior is fairly dispersed, with 36.6% of the distribution falling below 1. Values below 1 correspond to a scenario where GPC/LSC laboratories have smaller mean offsets than those mean offsets associated with AMS laboratories.

The modelled laboratory offset parameters, C_0 [lab i], have mean posterior parameter values ranging from -29.0 to 26.3 ¹⁴C years across the 68 laboratories (Table S3.1; Figure S3.4a). These posteriors show high overlap. At first glance, laboratory offset posteriors appear to show that between-laboratory variability is much higher than within-laboratory variability (Figure S3.4). However, this is only in the hypothetical scenario where a laboratory reports 0 measurement error. When measurement error is included, modelled within laboratory σ values increase rapidly and exceed the mean laboratory offsets (Figure S3.5; Figure S3.3e).

Posterior distributions for σ parameters are expressed on the log scale (Figure S3.5c–e). The posterior for the global parameter for σ , σ_{l} , has a mean value of 2.8 when this distribution is exponentiated and transformed back into the scale of ¹⁴C years (95% HPDI: 0.7–5.6). This represents within-laboratory measurement repeatability for the average AMS laboratory when reported ¹⁴C measurement error is zero. When σ_{lab} distributions are added to this average value, within-laboratory repeatability varies between laboratories. For the 68 labs in this dataset, the mean posterior σ value ranges from 1.53 to 9.92 ¹⁴C years (Table S3.1). However, these posterior distributions are dispersed and show considerable overlap. In general, the effect of σ_{AMS} causes GPC/LSC laboratories to have higher within-laboratory variability than AMS laboratories (Figure S3.3d; Figure S3.4b; Figure S3.5).

The generative aspect of this model allows one to simulate hypothetical pairs of laboratory parameters (Figure S3.6). As expected, GPC/LSC laboratories have generally larger mean offsets and within-laboratory σ values than AMS laboratories.



Figure S3.3. Posterior densities, posterior means (dots), and 95% highest posterior density intervals (HPDI) for model parameters. (a) Posteriors for each median-centered sample material. The sample medians are displayed to the right of each density to indicate over which ¹⁴C years the non-centered distributions fall. (b) Posterior distribution for C_{AMS} , which adjusts mean laboratory offsets if they are GPC/LSC measurements. Note, most of the density sits above 1, indicating that GPC/LSC laboratory offsets are probably more dispersed than their AMS counterparts. (c–d) Posterior distributions for the standard deviation parameters (σ_{I} , σ_{AMS} , and σ_{ME} , which are displayed on the log scale. (f–g) Posterior distributions for the standard deviations for the standard deviations of the distributions of each laboratory specific parameter, σ_0^{σ} and σ_{Iab}^{σ} (i.e., mean laboratory-specific standard deviations).



Figure S3.4. Posterior distributions for (a) mean laboratory offsets and (b) within-laboratory standard deviations. Gold densities show AMS laboratories and purple densities show GPC/LSC laboratories. GPC/LSC effects on laboratory offsets and within laboratory standard deviation values are included in these posterior distributions. Black dots mark the medians of each distribution. Note, within-laboratory standard deviations (b) assume 0 reported measurement error, and, in practice, these values become larger (see Figure S3.5). X-axes are on the ¹⁴C year scale.



Figure S3.5. Three examples showing how reported measurement error (ME) affects within-laboratory standard deviations: (a) ME = 25 ¹⁴C years, (b) ME = 50 ¹⁴C years, and (c) ME = 100 ¹⁴C years. Vertical blue lines mark the reported error values. Gold posterior densities are AMS laboratories and purple posterior densities are GPC/LSC laboratories. Black dots mark median values in each posterior density. GPC/LSC effects on within-laboratory standard deviations are included in these posterior distributions. X-axes are on the ¹⁴C year scale.



Figure S3.6. Laboratory parameters for (a) AMS and (b) GPC/LSC laboratories. Blue dots show 1000 simulated laboratories and red dots show mean posterior values for laboratories in the VIRI dataset. Both axes are displayed on the ¹⁴C year scale.

Table S3.1. Posterior means and 95% highest posterior density intervals (HPDI) for VIRI laboratory parameters. The effects of GPC/LSC measurement methods are excluded here. In other words, all laboratories are treated here as AMS laboratories to express variability that is due to laboratory identity exclusive of ¹⁴C measurement method.

		ry mean offset (¹⁴ C		idard deviation (¹⁴ C years)
Lab. ID [j]	yea	nrs) Co[lab j]	exp(o	_{ab} [lab j] + σι)
1	1.27	(-48.45 – 47.78)	3.54	(0.41 – 8.77)
2	14.87	(-38.29 – 71.87)	7.89	(1.44 – 17.99)
3	21.31	(-8.09 – 56.94)	2.88	(0.65 – 6.11)
4	-14.20	(-42.09 – 13.91)	2.32	(0.47 – 5.08)
5	-7.52	(-29.32 – 14.78)	2.09	(0.47 – 4.42)
6	11.97	(-18.54 – 41.52)	2.62	(0.41 – 6.40)
7	7.09	(-31.19 – 49.82)	2.85	(0.48 – 6.37)
8	11.62	(-25.75 – 47.68)	4.54	(0.94 – 9.87)
9	8.95	(-30.30 – 49.86)	2.18	(0.24 – 5.52)
10	9.47	(-27.11 – 45.79)	2.32	(0.25 – 5.59)
12	21.16	(-27.98 – 77.48)	9.92	(2.11 – 21.35)
13	5.61	(-31.05 – 41.74)	5.30	(1.25 – 11.18)
14	-12.61	(-65.43 – 37.58)	3.70	(0.53 – 8.91)
15	-16.92	(-52.35 – 17.87)	2.44	(0.42 – 5.30)
16	7.74	(-10.95 – 26.83)	1.82	(0.35 – 3.92)
17	-1.70	(-51.05 – 46.33)	3.49	(0.41 – 8.72)
20	-8.79	(-44.27 – 26.74)	2.24	(0.19 – 5.50)
21	11.63	(-24.49 – 43.44)	2.19	(0.18 – 5.57)
22	18.04	(-6.20 – 42.05)	3.03	(0.76 – 6.40)
23	-0.60	(-23.73 – 22.72)	1.53	(0.24 – 3.42)
24	2.00	(-38.73 – 45.11)	2.21	(0.20 – 5.35)
25	7.93	(-4.37 – 20.12)	1.60	(0.49 – 3.10)
26	21.80	(-8.74 – 50.63)	2.65	(0.58 – 5.81)
27	10.90	(-17.98 – 41.21)	1.92	(0.37 – 4.29)

28	-9.63	(-58.45 – 39.77)	4.14	(0.78 – 9.69)
29	-6.62	(-55.05 – 43.62)	2.42	(0.22 – 6.21)
30	9.20	(-35.61 – 55.32)	4.28	(0.69 – 9.66)
31	11.28	(-33.59 – 57.48)	4.85	(0.84 – 10.67)
32	-18.65	(-73.86 – 29.13)	3.62	(0.47 – 8.45)
33	-15.23	(-59.45 – 28.70)	2.42	(0.27 – 6.03)
34	-21.21	(-73.84 – 29.64)	3.47	(0.30 – 8.31)
35	23.47	(-17.02 – 63.13)	3.03	(0.33 – 7.25)
36	-10.87	(-54.42 - 29.74)	2.47	(0.33 – 6.01)
37	-15.38	(-45.41 – 15.47)	4.73	(1.16 – 9.54)
39	-29.02	(-84.32 - 20.78)	5.36	(0.81 – 11.78)
42	-11.70	(-46.12 - 24.47)	1.80	(0.26 – 4.24)
43	-11.68	(-32.10 - 9.40)	2.02	(0.44 - 4.22)
44	13.21	(-20.36 – 44.39)	3.24	(0.59 – 7.06)
45	-25.20	(-52.19 – 2.26)	3.59	(0.82 – 7.41)
46	19.72	(-0.37 – 39.30)	1.95	(0.45 – 4.03)
47	26.32	(-1.93 – 52.93)	2.53	(0.53 – 5.38)
48	-8.16	(-59.22 – 40.86)	5.93	(0.94 – 13.58)
50	-18.49	(-51.68 – 12.90)	1.58	(0.20 – 3.61)
53	2.49	(-21.70 – 24.94)	1.82	(0.34 – 3.93)
54	-8.95	(-48.71 – 30.48)	2.26	(0.25 – 5.58)
55	-5.29	(-34.24 – 22.08)	1.67	(0.14 – 4.15)
56	9.84	(-21.37 – 39.56)	1.79	(0.30 – 4.04)
58	-3.11	(-32.11 – 28.43)	2.14	(0.17 – 5.16)
59	11.95	(-35.21 – 61.90)	2.84	(0.29 – 6.99)
60	-26.96	(-53.13 – 0.31)	2.10	(0.37 – 4.65)
61	-9.55	(-58.01 – 39.37)	4.40	(0.62 – 10.24)
62	-6.03	(-38.43 – 26.03)	6.26	(1.31 – 13.27)
63	-14.42	(-60.64 – 27.33)	5.83	(1.07 – 12.76)
64	-28.35	(-60.61 – 6.38)	2.55	(0.44 – 5.71)
65		(-29.71 – 45.51)	2.62	(0.35 – 6.23)
66	13.15	(-21.49 – 52.30)	3.26	(0.54 – 7.00)
67	10.71	(-34.82 - 57.40)	3.77	(0.62 - 8.79)
69	13.29	(-30.11 - 60.55)	3.29	(0.50 – 7.46)
70	15.81	(-17.36 – 51.36)	3.53	(0.74 – 7.60)
71	-11.08	(-55.76 - 34.09)	2.24	(0.22 – 5.64)
72	-12.08	(-65.29 – 41.20)	4.21	(0.56 – 10.05)
73	-22.75	(-50.96 - 6.54)	1.92	(0.28 – 4.38)
74	12.59	(-38.32 - 67.89)	4.98	(0.68 – 11.78)
76	23.24	(0.83 – 44.15)	1.98	(0.34 – 4.50)
70	-3.51	(-51.84 - 47.22)	3.24	(0.34 – 4.30)
78	-13.89	(-60.91 – 29.29)	4.11	(0.60 - 9.37)
70 79	-6.77	(-42.76 - 29.40)	2.61	(0.32 - 6.37)
79 82	-0.77	(-33.75 - 69.33)	3.67	(0.32 - 0.37) (0.40 - 9.18)
02	10.92	(-00.10 - 09.00)	5:07	(0.70 - 0.10)

3.4. HMC Diagnostics and Posterior Predictive Checks

To ensure convergence of HMC chains, we examined trace plots, R-hat values, and effective sample sizes. Trace plots indicate convergence of HMC chains for all parameters. All R-hat values are below 1.01. Effective sample sizes close to 10,000 indicate efficient sampling of posterior parameter spaces. Of the 148 parameters, 101 have effective sample sizes of 10,000. The median effective sample size is also 10,000, the mean is 8362, and the minimum is 1056.4 (Figure S3.7).

We also performed a posterior predictive check by simulating data from the model and plotting these simulated data against the VIRI ¹⁴C measurements (Figure S3.8). For simplicity, where a laboratory contributed multiple ¹⁴C measurements for a given sample material, we used the average reported measurement error for that laboratory ID and sample (for example, laboratories 5 and 22 in the panel for Sample G). In all but nine of the 361 VIRI ¹⁴C measurements (97.5%), the 95% prediction intervals overlap the observed values. This indicates that the LBM does a reasonable job recovering the observed values.



Figure S3.7. Distribution of effective sample sizes across parameters. Parenthetical values show the number of parameter distributions that were sampled for each parameter type.



Figure S3.8. Posterior predictive check for the LBM. Each panel is a sample material reported in the VIRI study (Scott et al. 2010a,b; Scott, Cook, Naysmith et al., 2007). Laboratory IDs are listed along the y-axes. Reported ¹⁴C measurements and their associated errors (1 σ) are indicated by vertical and horizontal black segments, respectively. Horizontal gold and purple bars show the 95% posterior prediction intervals for AMS and GPC/LSC laboratories, respectively. The dashed red lines and bands indicate the mean and 95% highest posterior density intervals for the ¹⁴C value of each sample material.

4. Site and ¹⁴C Sample Selection for the YDB and LST

Our simulations included only those ¹⁴C measurements that are associated with samples from materials located within the YDB or LST stratigraphic layers. In this section, we list the sites with ¹⁴C measurements taken on samples from these layers and describe those samples. Where applicable, we detail why certain sites or samples were excluded from the simulations.

4.1. YDB Sites and ¹⁴C Samples

We compiled a preliminary list of ¹⁴C measurements for the YDB based on those samples that Kennett et al. (2015) present as originating from within a YDB layer. From there, we assessed each primary source to ensure the accuracy of each measurement and its provenience. The table below outlines the decisions made for each sample. Two samples that Kennett et al. (2015) identified as originating from the YDB, yet were excluded in our simulations, are described in rows highlighted in orange. Twenty-one rows highlighted in blue describe samples that other publications describe as associated with a YDB layer while Kennett et al. (2015) indicate otherwise. We did not include these samples in our simulations.

Lab Number	Reporte	ed ¹⁴C	Material	References		
μα		σ				
Abu Hureyra, Syri	a		L			
UCIAMS-105429	11,070	40	Charcoal	Bunch et al., 2012; Wittke et al., 2013		
OxA-172	10,900	200	Charred Seed	Bunch et al., 2012; Moore et al., 2000; Wittke et al., 2013		
OxA-430	11,020	150	Charred bone	Bunch et al., 2012; Moore et al., 2000; Wittke et al., 2013		
OxA-468	11,090	150	Charred bone	Bunch et al., 2012; Moore et al., 2000; Wittke et al., 2013		
BM-1718R	11,140	140	Charcoal	Bunch et al., 2012; Moore et al., 2000; Wittke et al., 2013		
Discussion						
YDB layer. Sample OxA-172 (Moore et al., 2000) was originally interpreted by Wittke et al. (2013) as being located within the YDB layer at Abu Hureyra. However, Kennett et al. (2015, SI8) state that OxA- 172 is "adjacent" to sample UCIAMS-105429, but not within the YDB layer. We default to Kennett et al. (2015) regarding the provenience of these samples and have included sample UCIAMS-105429 but not sample OxA-172 in our simulations. Bunch et al. (2012) identified samples OxA-430, OxA-468, and BM-1718R as close to the YDB layer via an age-depth model, although it is not clear that these samples were located within the YDB layer. Given their unsecure spatial relationship to the Abu Hureyra YDB layer, we deferred to Wittke et al. (2013) and Kennett et al. (2015) and excluded these samples from our simulations.						
Arlington Canyon			d States			
UCIAMS-47239	11,105	30	Charcoal	Kennett et al., 2008		
UCIAMS-36308	11,095	25	Wood	Kennett et al., 2008		
UCIAMS-42816	11,095	25	Wood	Kennett et al., 2008		
UCIAMS-36307	11,070	25	Wood	Kennett et al., 2008		
UCIAMS-36961	11,440	90	Carbon elongate	Kennett et al., 2008		
UCIAMS-36960	11,185	30	Carbon spherule	Kennett et al., 2008		
UCIAMS-36962	11,110	35	Wood	Kennett et al., 2008		
UCIAMS-36959	11,075	30	Glassy carbon	Kennett et al., 2008		
Beta-161032	10,860	70	Charcoal	Kennett et al., 2008		
UCIAMS-36306	11,375	25	Wood	Kennett et al., 2008		

UCIAMS-36305	44.005	25	Wood	Kennett et al. 2000			
UCIAMS-36304	11,235 11,105	25 30	Wood	Kennett et al., 2008 Kennett et al., 2008			
Discussion	11,105	30	WOOd	Refinell et al., 2006			
	(F) indicat	a that a	1 12 of these same	es are associated with the YDB at Arlington			
			asurements in our sir				
Barber Creek, Nor							
				ed from the YDB layer.			
Big Eddy, Missou			on samples recover				
AA-27486	11,900	80	Charcoal	Hajic et al., 2007			
AA-27480 AA-26654	10,710	85	Charcoal	Lopinot et al., 1998; Wittke et al., 2013			
AA-20034 AA-25778	10,710	85	Wood charcoal	Hajic et al., 2007			
AA-25776 AA-72612	10,200	54	Charcoal	Lopinot et al., 2000; Wittke et al., 2013			
Discussion	10,959	-04	Charcoar				
Wittke et al. (2013) Kennett et al. (201 within the same stra from an additional f (2015) rejected AA excluded it from ou Wittke et al. (2013)	identified 5) to differ atum. In co four sample A-25778 as ir simulatio identify A4	a peak i rentiate = ontrast, H es within s an out ons. A-72612	n nanodiamonds bet samples within these lajic et al. (Hajic et a that stratum. Althou lier based on an Ox as associated with th	ated with the YDB, AA-27486 and AA-26654. ween 327 and 335 cm below surface, leading e depths from samples taken at other depths I., 2007) do not distinguish these two samples gh located within the YDB layer, Kennett et al. (Cal age-sequence model and we have thus the YDB, although Kennett et al. (2015) indicate e the YDB. We excluded AA-72612 from our			
simulations.							
simulations. Blackwater Draw, New Mexico, United States							
Blackwater Draw,	New Mex	ico, Uni	ieu Sidies				
SMU-1880	New Mex 10,780	110 110	Soil humate	Johnson and Holliday, 1997; Wittke et al., 2013			
SMU-1880 Discussion	10,780	110	Soil humate	2013			
SMU-1880 Discussion Wittke et al. (2013 (measurement orig include this sampl indicates that this s potential YDB sam SMU-1880 as asso	10,780 3) report f inally repo e in their sample was ples are re ociated with	110 that san rted by s narrative s incorpo ported in the YD	Soil humate nple SMU-1880 is I Johnson & Holliday (e description, figures prated into an age-se n the literature. Since B, we do not include				
SMU-1880 Discussion Wittke et al. (2013 (measurement orig include this sampl indicates that this s potential YDB sam SMU-1880 as asso Blackville, South	10,780 3) report f inally repo e in their sample was ples are re poiated with Carolina,	110 that san rted by s narrative s incorpo s incorpo ported in the YD United s	Soil humate nple SMU-1880 is I Johnson & Holliday (e description, figures prated into an age-se n the literature. Since B, we do not include States	2013 ocated within the YDB at Blackwater Draw 1997)). However, Kennett et al. (2015) do not s, or tables. Their supplemental OxCal code equence model as a potential outlier. No other e Kennett et al. (2015) do not explicitly identify this measurement in our simulations.			
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GrA-49574	10,845	45	Charcoal	van Hoesel et al., 2012
GrA-49569	10,895	45	Charcoal	van Hoesel et al., 2012
GrA-49509 GrA-49514	10,880	110	Charcoal	van Hoesel et al., 2012
GrA-49575	10,880	50	Charcoal	van Hoesel et al., 2012
Discussion	10,900	50	Charcoar	Vall HUESELELAL, 2012
	(2012)	ant 1 1 14		wanta from the charged rich llocale llorizon
				nents from the charcoal-rich Usselo Horizon,
				tential to evaluate the Younger Dryas Impact
				ample measurements on the basis that they hich does not have nanodiamond markers. In
				ced these three YDB samples stratigraphically
				only included ¹⁴ C measurements from these
three samples.	iy i i saili	pies. In	our simulations, we	only included to measurements nom these
Indian Creek, Mor	ntana Unit	ted Stat	06	
				ed from the YDB layer.
Lake Cuitzeo, Mic			on samples recover	
				ed from the YDB layer.
			on samples recover	eu lioni line f DB layer.
Lake Hind, Manito			Characal	Eirestens et al. 2007
UCIAMS-29317	10,610	25	Charcoal	Firestone et al., 2007
Discussion	4 5) in line		La dia da MDD 140	
				sample is located below a peat layer, citing
				do not describe the sample used for the ¹⁴ C
				nats, microspherules, glass-like carbon, and
				describe the context of the sample.
Lindenmeier, Colo	brado, Uni	ted Sta	tes	
I-141	10,780	135	Charcoal	Haynes and Agogino, 1960; Kinzie et al.,
	-,		-	2014; Walton et al., 1961
Discussion			11.0	
				obtained from the nanodiamond rich layer that
				stratigraphically directly above the YDB (Kinzie
			I. (2015), who indica	te that this sample dates the YDB layer, and
have included it in	our simula	tions.		
1 4 4 4 finat much	liabad an A	0 700 .		and America, 1000) and later some stad to
				es and Agogino, 1960) and later corrected to
			1961). Like Kennett	et al. (2015), we use the corrected error term.
Lingen, Lower Sa				1/
Beta-369246	10,870	40	Charcoal	Kennett et al., 2015
Discussion				
				by an abundance of charcoal combined with
				al. (2015) rely on this identification for their
association of this	<u> </u>	h the YL	DB.	
Lommel, Belgium				
UCIAMS-46303	11,480	100	Charcoal	Wittke et al., 2013
N/A	10,950	50	N/A	van Geel et al., 1989; Wittke et al., 2013
Discussion				
				m within the YDB, conflicting with a previous
				nent on charcoal from the Lommel YDB layer
				t on charcoal of 10,950±50 ¹⁴ C yr BP, although
				et al., 2013). Wittke et al. (2013) cite van Geel
				ever, van Geel et al. (1989) appear to have
				s rather than report a ¹⁴ C measurement on a
			why this value is no	t associated with a provenience or laboratory
ID, it is not a ¹⁴ C m	easureme	nt.		
				tigraphically below the YDB, and they do not
include the latter ¹	⁴ C estimat	te. No e	explanation is provid	ed for these discrepancies with Wittke et al.

(2013) Due to the	uncertain	nrovenia	ance of these ¹⁴ C sa	mples, we followed Kennett et al. (2015), and
			ne Lommel YDB in ou	
Melrose, Pennsylv				
				ed from the YDB layer.
Mucuñuque, Vene		rementa		
		romonte	on complex recover	ed from the YDB layer.
Murray Springs, A				
A-1045				Havaa 2007: Wittka at al. 2012
	10,760	100	Charcoal + F ₂	Haynes, 2007; Wittke et al., 2013
TX-1045	10,260	140	Humates + F ₂	Haynes, 2007
TX-1044	12,600	2440	Charcoal + F ₂	Haynes, 2007
TX-1462 Discussion	10,930	170	Charcoal	Haynes, 2007; Wittke et al., 2013
entirely from the na 1462 is apparently i despite being abser did not include TX- Springs ¹⁴ C measu 26212, A-1045, an analyses. We deferred to Ken 1462. Given the en	arrative, fig included in nt elsewhe -1045 or T rements th id TX-1462 anett et al.'s xtreme ¹⁴ C ted in our	their Ox re in the X-1044 at are ir 2). We s (2015) c measu main p	nd tables in Kennett Cal age-sequence n e text. Further confusi in their age estimat ncorrectly listed as O suspect that this is list of YDB ¹⁴ C samp urement error for TX aper. We did, howe	52 dates the YDB, but this sample is missing et al. (2015). An additional conflict is that TX- nodel for Murray Springs (Kennett et al., 2015), ion arises from the fact that Wittke et al. (2013) ions of the YDB. Of note are several Murray SL ages in Wittke et al. (2013): Table S1 (AA- a typographic error that did not impact their oles for Murray Springs, thereby excluding TX- (-1044, we also excluded this sample in the ver, include this sample in simulations of an
Ommen, Netherla		11 0000		
UCIAMS-46307	11,440	35	Charcoal	Wittke et al., 2013
Discussion	,			
with Wittke et al. (charcoal in the Y stratigraphically be include any sample	(2013), wh (DB. Kenr low the YI es from Om	io repor nett et DB. For	t that AMS sample al. (2015) report t	covered from the Ommen YDB. This conflicts UCIAMS-46307 was recovered directly from hat this sample originates from a context deferred to Kennett et al. (2015) and did not
Santa Maira, Spain		1.10		A T () 0000
Beta-75225	11,020	140	Charcoal	Aura Tortosa et al., 2008
Discussion				
				sample for the Santa Maira YDB.
Sheriden Cave, Ol				
UCI-38249-(C)	10,915	30	Bone Clovis point	Waters et al., 2009
Beta-127909	10,840	80	Wood charcoal	Tankersley and Redmond, 1999
Beta-127910	10,960	60	Wood charcoal	Tankersley and Redmond, 1999
Discussion				
				re associated with the YDB at Sheriden Cave, luded all 3 measurements in our simulations.
Talega, California	, United S	tates		
Beta-196150	11,070	50	Charcoal	Bergin, 2011 in Kennett et al., 2015; Wittke et al., 2013
Discussion				
We followed Kenne	ett et al. (20	015) and	d included this single	sample for the Talega YDB.
Topper, South Ca				
. oppol, ooutil ou	rolina, Un	ited Sta	ites	
AA-100294	10,958 10,958	65	Charcoal	Goodyear, 2013

4.2. LST Sites and ¹⁴C Samples

Many of the samples described here were originally summarized in Baales et al. (2002). Three rows highlighted in orange describe ¹⁴C samples recovered from within the LST and summarized by Baales et al. (2002) yet excluded in our simulations for reasons specified in the associated discussion paragraphs. Ten rows highlighted in blue describe ¹⁴C samples recovered from stratigraphic contexts near the LST, but not from within it. Since these samples likely do not date the Laacher See volcanic eruption, we excluded them from our simulations.

Laboratory			Material	Original Reference
Number				
Brohl Valley				
HV-11774	11,075	185	Plant remains	Heine, 1993
HD-17900	11,277	26	Tree 1/4, rings 1–38	Kromer et al., 1998
KN-3800	11,240	100	Populus	Street, 1993
KN-3801	11,260	95	Populus	Street, 1993
KN-3802	11,280	100	Populus	Street, 1993
KN-3803	11,510	90	Populus	Street, 1993
Unknown	11,085	90	Charcoal	Frechen, 1952; Schweitzer, 1958
HD-17100	11,206	20	1a ca. 50 rings	Kromer et al., 1998
HD-17145	11,223	22	3a ca. 50 rings	Kromer et al., 1998
HD-17101	11,121	28	5b ca. 50 rings	Kromer et al., 1998
Discussion			<u>_</u>	
Neither Frechen (19	52) nor Sch	weitzer (1	958) report the laboratory number	er for the unknown sample.
Glees				
GrA-?	10,680	85	Charcoal	Frechen, 1959;
GIA-:	10,000	00	Charcoar	Schweitzer, 1958
Discussion				
Originally reported number.	by Schweitz	er (1958)). Frechen (1959) provides the l	aboratory but not the sample
Kruft				
HD-19098	11,063	30	Populus 9 rings 1-20	Baales et al., 1998;
11D-19090	11,005	- 50	Fopulus 9 migs 1-20	Kromer et al., 1998
HD-18438	11,065	22	Populus 8 outer rings	Baales et al., 1998;
110-10-00	11,000	~~~		Kromer et al., 1998
HD-19092	11,066	28	Populus 9 rings 21-30	Baales et al., 1998;
110 10002	11,000	20		Kromer et al., 1998
HD-18622	11,073	33	Populus 9 rings 31-40	Baales et al., 1998;
	,0.10			Kromer et al., 1998
HD-19037	11,075	28	Populus 9 rings 41-50	Baales et al., 1998; Kromer et al., 1998
HD-18648	11,037	27	Populus 1 rings 31-40	Baales et al., 1998; Kromer et al., 1998
Discussion	1	L	I	
	rements (HI	D-19098	HD-19092, HD-18622, and HD-1	9037) originate from the same
			/ included the sample correspond	
			corresponds to the only Populu	
			acher See eruption. The other thr	
eruption it in the ord				
Miesenhein IV				
OxA-3584	11,190	90	Alces alces bone	Hedges et al., 1993

for moss accumulation prior to deposition of the LST. While not enough time elapsed to fully surround the remains in the pre-LST layer, the event must necessarily have followed the death of the animal, and we have therefore excluded these measurements. Nette Valley W-525 10,800 300 Charcoal Frechen, 1959 N/A 10,880 95 Charcoal Van den Bogaard and Schmincke, 1985 Discussion In their supplementary data, van den Bogaard & Schmincke (1985) note that the second ¹⁴ C measurement is from a personal communication with Geyh in 1976. Soppensee ETH-5290 10,760 80 Macrofossils Hajdas et al., 1993, 1995 ETH-6930 11,190 80 Macrofossils Hajdas et al., 1993, 1995 ETH-12617 11,040 90 Macrofossils and wood/bark Hajdas et al., 1995 ETH-12615 11,370 90 Macrofossils and wood/bark Hajdas et al., 1995 ETH-12613 11,220 90 Macrofossils and wood/bark Hajdas et al., 1995 ETH-12610 11,180 100 Macrofossils and wood/bark Hajdas et al., 1995 Discussion 14,220 90 Macrofossils and wood/bark Hajdas et al., 1995	OxA-3585	11,310	95	Alces alces bone	Hedges et al., 1993					
Hedges et al. (1993) note that the Alces remains predate the Laacher See eruption as there was time for moss accumulation prior to deposition of the LST. While not enough time elapsed to fully surround the remains in the pre-LST layer, the event must necessarily have followed the death of the animal, and we have therefore excluded these measurements. Netzet Valley W-525 10,800 300 Charcoal Frechen, 1959 N/A 10,880 95 Charcoal Schmincke, 1985 Discussion In their supplementary data, van den Bogaard & Schmincke (1985) note that the second ¹⁴ C measurement is from a personal communication with Geyh in 1976. Soppensee ETH-5290 10,760 80 Macrofossils Hajdas et al., 1993, 1995 ETH-6930 11,190 80 Macrofossils Hajdas et al., 1993, 1995 ETH-6932 10,540 150 Macrofossils and wood/bark Hajdas et al., 1993, 1995 ETH-12615 11,370 90 Macrofossils and wood/bark Hajdas et al., 1995 ETH-12613 11,220 90 Macrofossils and wood/bark Hajdas et al., 1995 Discussion Hajdas et al., 1995 Intervent the sample or if the sample or ginate from within ash. Hajdas et al., 1995 Hajdas et al. (1995) report the bottom four ¹⁴ C measurements on samples from 1-2 cm sediment slicesthat also contain	OxA-3586	11,190	100	Alces alces bone	Hedges et al., 1993					
for moss accumulation prior to deposition of the LST. While not enough time elapsed to fully surround the remains in the pre-LST layer, the event must necessarily have followed the death of the animal, and we have therefore excluded these measurements. Nette Valley W-525 10,800 300 Charcoal Frechen, 1959 N/A 10,880 95 Charcoal Van den Bogaard and Schmincke, 1985 Discussion In their supplementary data, van den Bogaard & Schmincke (1985) note that the second ¹⁴ C measurement is from a personal communication with Geyh in 1976. Soppensee ETH-5290 10,760 80 Macrofossils Hajdas et al., 1993, 1995 ETH-6930 11,190 80 Macrofossils and wood/bark Hajdas et al., 1993, 1995 ETH-12617 11,040 90 Macrofossils and wood/bark Hajdas et al., 1995 ETH-12615 11,370 90 Macrofossils and wood/bark Hajdas et al., 1995 ETH-12610 11,180 100 Macrofossils and wood/bark Hajdas et al., 1995 Discussion Hajdas et al., 1995 Intelias at al. (1995) report the bottom four ¹⁴ C measurements on samples from 1-2 cm sediment sliceer that also contain ash from the Laacher See eruption. However, the relationship between the sample materials and the ash within each slice is unkno	Discussion									
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Tönnisstein Rubin and Alexander,		ides the lab	oratorv bi	it not the sample number.						
W 528 11 150 200 Charcoal Rubin and Alexander,	· · · · ·									
		11,150	200	Charcoal						
GrA-? 11,025 90 Charcoal Frechen, 1959	GrA-?	11,025	90	Charcoal						
Discussion	Discussion									
Frechen (1959) provides the laboratory but not the sample number for the latter sample.		ides the lab	oratory bu	ut not the sample number for the l	atter sample.					

5. Simulations of Alternative LST_{Obs} and YDB_{Obs}

In total, this "multiverse analysis" entailed 42 separate simulations over six different parameterizations (A1–C1 and A2–C2) and across seven different observed datasets (the main text dataset scored as all AMS measurements, as well as the three alternative datasets, each run both with the reported AMS/GPC/LSC distinctions and with all measurements scored as AMS). When describing the results for each alternative, we focus on simulation C2, much like the results described in the main paper. We focus on simulation C2 because it contains the most sources of variability, which likely most closely approximates the large number of sources that must affect real ¹⁴C datasets (i.e., C2 is characterized the most realism of the six simulations). The remaining five simulations show expectations when these sources of variability are excluded (or their effects are relaxed, as is the case for the OWM simulations with λ set to 0.04).

5.1. Alternative 1

The Alternative 1 dataset includes the five measurements that were excluded from simulations in the main text. The σ^{14} C and dissimilarity values for both YDB_{Obs} and LST_{Obs} become a more probable outcome across simulations for Alternative 1 (Figure S5.1). However, this change is modest for YDB_{Obs}, which remains a highly improbable event, with less than 0.1% of simulated σ^{14} C and dissimilarity values exceeding the observed σ^{14} C and dissimilarity values in any given simulation.

Treating all Alternative 1 measurements as AMS decreases the probability of observing a dataset with $\sigma^{14}C$ and dissimilarity values as large as those in either observed dataset. This decrease is largest for LST_{Obs}, which is dominated by non-AMS measurements. Despite the decrease having a disproportionately large effect on the probability of observing $\sigma^{14}C$ and dissimilarity values for LST_{Obs}, the dispersion statistics associated with LST_{Obs} remain orders of magnitude more probable than those associated with YDB_{Obs}.

5.2. Alternative 2

Alternative 2 excludes three YDB_{Obs} measurements with potential reliability issues, which may have artificially increased dispersion in YDB_{Obs}. LST_{Obs} retains all 19 measurements used in the main text for this event. This alternative was designed to be favorable to the Younger Dryas Impact Hypothesis by reducing dispersion in the observed dataset ¹⁴C measurements.

While simulated datasets as dispersed as YDB_{Obs} do become more probable when these three measurements are excluded, YDB_{Obs} σ^{14} C and dissimilarity values remain highly improbable, with less than 0.5% of iterations producing σ^{14} C and dissimilarity values greater than those values for YDB_{Obs} in any simulation (Figure S5.1). The minor LST results differences between the main text and the results presented here are due to simulation variance, given that both simulations used the same set of LST_{Obs} measurements.

Much like Alternative 1, scoring all Alternative 2 measurements as AMS reduces the probability of observing σ^{14} C and dissimilarity values as large as those in either observed dataset. However, this reduction is slight for YDB_{Obs}. Although simulated LST σ^{14} C and dissimilarity values are more greatly affected by scoring all observed measurements as AMS, the LST_{Obs} σ^{14} C and dissimilarity values remain probable relative to the YDB_{Obs} values, in the context of the simulation.

5.3. Alternative 3

The Alternative 3 dataset excludes 22 measurements from sites that make up a disproportionately large share of YDB_{Obs} and LST_{Obs}. For YDB_{Obs}, these excluded measurements

are the samples from Arlington Canyon (n = 12), and for LST_{Obs}, these measurements are the samples from Brohl Valley (n = 10).

YDB_{Sim} σ^{14} C and dissimilarity values exceed YDB_{Obs} σ^{14} C and dissimilarity values more often for Alternative 3 than they do for the observed dataset used in the main text, but this gain is slight (Figure S5.1). Even when these Arlington Canyon samples are excluded, simulations produced σ^{14} C and dissimilarity values exceeding those values for YDB_{Obs} in no more than 0.1% of iterations. In contrast, when the Brohl Valley measurements are excluded from LST_{Obs}, LST_{Sim} σ^{14} C and dissimilarity values more often exceed LST_{Obs} σ^{14} C and dissimilarity values. In those simulations that include the LBM, this ranges from 7.5–28.3% of simulated σ^{14} C values and 10.0– 43.6% of dissimilarity values.

Scoring all Alternative 3 measurements as AMS reduces the number of YDB_{Sim} and LST_{Sim} datasets with σ^{14} C and dissimilarity values that exceed those values for YDB_{Obs} and LST_{Obs}. This reduction is greatest for LST measurements in Alternative 3, but even with this larger reduction, the LST_{Obs} σ^{14} C and dissimilarity values remain much more probable than do the YDB_{Obs} σ^{14} C and dissimilarity values remain for those values.

5.4. Main text LST_{Obs} and YDB_{Obs} measurements scored as entirely AMS

As with the alternative datasets, we also completed the simulation for the main text dataset with all measurements scored as AMS (Figure S5.1). YDB_{Sim} σ^{14} C values exceed YDB_{Obs} σ^{14} C values slightly more often when all measurements are scored as AMS, while simulated dissimilarity values exceed observed dissimilarity values *less* often. In all cases, the all-AMS YDB simulations produce σ^{14} C or dissimilarity values that exceed the observed values in no more than 0.1% of iterations. The number of LST_{Sim} iterations that produced σ^{14} C or dissimilarity values exceeding the LST_{Obs} σ^{14} C and dissimilarity values decreased as a result of scoring all measurements as AMS, but even with this decrease, the LST_{Obs} σ^{14} C and dissimilarity values remain much more probable in the context of the simulation than do the YDB_{Obs} σ^{14} C and dissimilarity values.

5.5. Discussion

The alternative simulations demonstrate that inferences from our simulations vary with data inclusion decisions, although the variation that we investigated here does not substantially alter the qualitative inferences presented in the main text: The amount of dispersion in the YDB_{Obs} ¹⁴C measurements is highly improbable, given a synchronous event. In contrast, dispersion in the LST_{Obs} ¹⁴C measurements, while not highly probable, is orders of magnitude more probable than the dispersion in the YDB_{Obs} measurements, given a synchronous event. This difference is especially interesting for Alternative 2, in which the YDB_{Obs} dissimilarity value is only 0.01 greater than the LST_{Obs} dissimilarity value, and the YDB_{Obs} σ^{14} C value falls *below* the LST_{Obs} σ^{14} C value. This highlights the context-dependent aspect of these simulations-the degree of clustering in ¹⁴C measurements that should be expected depends on many variables associated with those measurements, including their reported measurement error, the number of possible "old wood" samples, the number of laboratories that contributed measurements, and the measurement methods employed by those laboratories. Although the observed σ^{14} C and dissimilarity values for Alternative 2 might suggest that the YDB_{Obs} measurements are more consistent with synchroneity than are the LST_{Obs} measurements, aspects of YDB_{Obs} strongly suggest that they should be much more clustered than this if these measurements are associated with a synchronous event.

Inferences also necessarily change given different choices about simulation design, which are theoretically infinite. We designed the simulations to include those variables that should have the largest effects on the dispersion of ¹⁴C measurements, but it is possible to imagine arguments for

other variables that we excluded, or arguments for different effects associated with the variables that we did include (e.g., the OWM entails choices about which value to specify for λ , as well as the choice to model "old wood" effects with an exponential distribution. See SI Section 6 for further discussion). We aimed to demonstrate a range of inferences given different simulation assumptions, but this range could be further extended with other simulation designs. An important observation for these simulations is that the expected dispersion within a set of synchronous ¹⁴C measurements generally declines when fewer sources of variability are considered. This is especially true of the laboratory variability described in the LBM. As such, we largely ignored the implications of the simulations that excluded the LBM, as it is unrealistic to expect a variety of laboratories with different protocols to measure ¹⁴C with perfect precision and accuracy. Such a scenario would necessarily underestimate the amount of dispersion expected in set of ¹⁴C measurements. This expectation is supported by multiple inter-laboratory ¹⁴C measurement studies (Boaretto et al., 2003; International Study Group, 1982; Scott et al., 1990, 1998, 2010a; Scott, Cook, Naysmith et al. 2007). We displayed the results from simulations that excluded the LBM to illustrate underestimation effects on σ^{14} C and dissimilarity values when this variability is ignored.

In addition to the inclusion or exclusion of the LBM, there is also the issue of choices in specifying the LBM (described in SI Section 3). Much like simulation design choices, the LBM would characterize inter- and intra-laboratory measurement variability differently under alternative specifications. We aimed to include those variables most likely to affect the dispersion of ¹⁴C values in this context, although other choices might plausibly be made. For example, rather than model the effect of AMS vs GPC/LSC measurements, one could model the effects of AMS vs GPC vs LSC measurements. We lumped the latter two measurement methods together since there are few data within each measurement method group when GPC/LSC measurements are subdivided into GPC and LSC measurements. A primary measurement distinction commonly drawn by researchers working with ¹⁴C datasets is that of AMS compared to earlier counting methods, such as GPC and LSC. As such, we incorporated this distinction into our model.



6. Simulations with Alternative Old Wood Model λ Values

In the main paper, simulations were run with old wood offsets drawn from exponential distributions with λ set to 0.04 and 0.01 (mean respective offsets, 25 and 100 years; 95% of respective expected offsets, 0–75 years and 0–300 years) as well as simulations that lacked old wood offsets. Reducing λ allows for larger offsets, increasing dispersion in simulated datasets. In the main paper, we selected λ values with the goal of bounding realistic old wood effects in each context of interest. However, λ values could theoretically be decreased until simulations generate datasets matching or exceeding dispersion in the observed datasets.

We ran additional simulations with reduced λ values to examine what degree of old wood age offsetting would be necessary to obtain the level of dispersion in YDB_{Obs}. These additional simulations set λ to 5.0e-3 and 2.5e-3. For the former λ value, the mean expected age offset is 200 years, with 95% of expected offsets falling within 0–600 years. For the latter λ value, the mean expected age offset is 400 years, with 95% of expected offsets falling within 0–1200 years. These offsets are generally large given the tree species identified in each dataset, and they likely overestimate old wood effects. For these alternative simulations, we used the observed datasets presented in the main paper rather than the various alternative datasets.

At $\lambda = 5.0e-3$, simulated dissimilarity and $\sigma^{14}C$ values indicate less dispersion in YDB_{Sim} datasets than in YDB_{Obs} (Figures S6.1 and S6.2). In contrast, LST_{Sim} are roughly as dispersed as LST_{Obs}, suggesting that this magnitude of old wood effects is consistent with a known synchronous event when the average old wood offset is at 200 calendar years. YDB_{Sim} datasets resemble or are more dispersed than YDB_{Obs} (Figures S6.3 and S6.4). This suggests that an average old wood offset between 200 and 400 years is required to produce dispersion in YDB_{Sim} datasets that is comparable to those in YDB_{Obs}.

The effects of old wood offsets begin to swamp those of the LBM as λ decreases, suggesting that laboratory error and repeatability has little explanatory power if YDB_{Obs} truly results from a synchronous event. When λ reaches 2.5e-3, dispersion in YDB_{Obs} falls within two z-scores of the mean dispersion in the simulated datasets, regardless of whether than LBM is included (Figures S6.3 and S6.4). As such, old wood effects must be a proportionally large source of variability in ¹⁴C measurements to account for the dispersion in YDB_{Obs}. We find it unrealistic to assume that the "old wood" measurements in this dataset precede the event of interest by an average of 400 years. Rather, a more conservative assumption, given the results, is that a single event was not responsible for YDB_{Obs}.





Figure S6.1. Results of simulations with λ = 5.0e-3 and no LBM. Blue geometry corresponds to LST simulations, and red geometry corresponds to YDB simulations. Refer to Fig. 4 (main text) for interpreting the segments and bands in each panel.

Figure S6.2. Results of simulations with λ = 5.0e-3 and the LBM. Blue geometry corresponds to LST simulations, and red geometry corresponds to YDB simulations. Refer to Fig. 4 (main text) for interpreting the segments and bands in each panel.





Figure S6.3. Results of simulations with λ = 2.5e-3 and no LBM. Blue geometry corresponds to LST simulations, and red geometry corresponds to YDB simulations. Refer to Fig. 4 (main text) for interpreting the segments and bands in each panel.

Figure S6.4. Results of simulations with λ = 2.5e-3 and the LBM. Blue geometry corresponds to LST simulations, and red geometry corresponds to YDB simulations. Refer to Fig. 4 (main text) for interpreting the segments and bands in each panel.

7. Sample Size Differences Between YDBobs and LSTObs

One feature that differs between YDB_{Obs} and LST_{Obs} is the number of measurements in each dataset. There are 30 measurements in YDB_{Obs} , while LST_{Obs} contains only 19 measurements. This raises the question of whether YDB_{Obs} could be more consistent with synchroneity if it was reduced to 19 measurements. Depending on which 19 measurements are subsampled from YDB_{Obs} , a reduced dataset of 19 measurements could be more or less dispersed than the full dataset of 30 measurements.

To understand the implications of differences in sample size, we ran the simulation using three subsampled datasets of 19 YDB_{Obs} measurements. To create these datasets, we first generated 10,000 randomly subsampled datasets from the 30 YDB_{Obs} measurements. We calculated σ^{14} C for each dataset, resulting in a distribution of 10,000 σ^{14} C values. Next, we selected the datasets associated with the 2.5%, 50.0%, and 97.5% percentile values in the σ^{14} C distribution. These represent plausible observed datasets of 19 measurements with low, medium, and high dispersion, given the known dataset of 30 YDB_{Obs} measurements (Tables S7.1 and S7.2).

The subsampled YDB_{Obs} dataset with low dispersion has a dissimilarity value comparable to the dissimilarity value for LST_{Obs}. The YDB_{Obs} σ^{14} C value is ~24.5 lower than the σ^{14} C value for LST_{Obs}, suggesting an observed YDB dataset that is potentially more consistent with synchroneity than is the observed LST dataset. Despite this, the observed statistics for the low dispersion subsample of YDB_{Obs} are still more dispersed than those expected from the simulations (Figure S7.1). In nearly all versions of the simulation, less than 0.1% of YDB_{Sim} dissimilarity and σ^{14} C values are greater than the dissimilarity and σ^{14} C values for YDB_{Obs}. The exceptions to this are the σ^{14} C values of the C simulations, which have the larger old wood offset ($\lambda = 0.01$). In simulation C1, which excludes the LBM, 1.6% of YDB_{Sim} σ^{14} C values exceed the YDB_{Obs} σ^{14} C value (Figure S7.1). In simulation C2, which includes the LBM, 2% of YDB_{Sim} σ^{14} C values exceed the YDB_{Obs} σ^{14} C value. Although these percentages are larger than those obtained from using the full 30 measurements in YDB_{Obs}, they are still low, especially compared to the LST simulations.

The subsampled YDB_{Obs} dataset with medium dispersion has dissimilarity and σ^{14} C values similar to those obtained for the YDB_{Obs} dataset with the full 30 measurements. As such, the simulated results resemble those presented for the full dataset in the main text (Figure S7.1). As expected, the subsampled YDB_{Obs} dataset with high dispersion is extremely inconsistent with the simulated datasets (Figure S7.1).

Results demonstrate that despite the sample size differences between LST_{Obs} (n = 19) and YDB_{Obs} (n = 30), it is highly improbable that a YDB_{Obs} dataset that has been rarefied to 19 measurements would be consistent with synchroneity. Although the low dispersion subsample is *more* consistent with synchroneity than is the full YDB_{Obs} dataset of 30 measurements, it remains a highly improbable outcome, given a synchronous event.

	Dispers	sion in subsampled YDB _{Obs} (datasets
	Low	Medium	High
σ ¹⁴ C	172.45	277.78	343.85
Dissimilarity	0.71	0.73	0.83

Table S7.1. Dispersion in subsampled YDB_{Obs} measurements.

			OWM	YDB	_{Obs} disp	
Site ID: ¹⁴ C Measurement	Material	AMS	sample	Low	Med.	Higł
G: 11,070 ± 40 (UCIAMS-105429)	Charcoal	1	1			Х
H: 11,105 ± 30 (UCIAMS-47239)	Charcoal	1	1	Х		Х
H: 11,095 ± 25 (UCIAMS-36308)	Wood	1	1		Х	Х
H: 11,095 ± 25 (UCIAMS-42816)	Wood	1	1	Х	Х	
H: 11,070 ± 25 (UCIAMS-36307)	Wood	1	1	Х	Х	
H: 11,440 ± 90 (UCIAMS-36961)	Carbon elongate	1	0		Х	Х
H: 11,185 ± 30 (UCIAMS-36960)	Carbon spherule	1	0	Х		Х
H: 11,110 ± 35 (UCIAMS-36962)	Wood	1	1	Х		Х
H: 11,075 ± 30 (UCIAMS-36959)	Glassy carbon	1	0	Х		
H: 10,860 ± 70 (Beta-161032)	Charcoal	1	1	Х	Х	
H: 11,375 ± 25 (UCIAMS-36306)	Wood	1	1			Х
H: 11,235 ± 25 (UCIAMS-36305)	Wood	1	1	Х	Х	Х
H: 11,020 ± 25 (UCIAMS-36304)	Wood	1	1			Х
K: 10,840 ± 75 (GrA-49524)	Charcoal	1	1	Х	Х	
K: 10,865 ± 55 (GrA-49509)	Charcoal	1	1	Х	Х	
K: 11,020 ± 75 (GrA-49515)	Charcoal	1	1		Х	
l: 11,900 ± 80 (AA-27486)	Charcoal	1	1		Х	Х
l: 10,710 ± 85 (AA-26654)	Charcoal	1	1	Х	Х	Х
J: 11,070 ± 60 (Beta-184854)	Sed. organics	1	0	Х	Х	Х
O: 10,760 ± 100 (A-1045)	Charcoal	0	1	Х	Х	Х
O: 10,260 ± 140 (TX-1045)	Humates	0	0			Х
Q: 10,915 ± 30 (UCI-38249)	Bone Clovis point	1	0	Х	Х	Х
Q: 10,840 ± 80 (Beta-127909)	Wood charcoal	1	1	Х	Х	Х
Q: 10,960 ± 60 (Beta-127910)	Wood charcoal	1	1	Х		
L: 10,610 ± 25 (UCIAMS-29317)	Charcoal	1	1	Х		Х
M: 10,780 ± 135 (I-141)	Charcoal	0	1	Х	Х	
N: 10,870 ± 40 (Beta-369246)	Charcoal	1	1		Х	
R: 11,070 ± 50 (Beta-196150)	Charcoal	1	1		Х	
S: 10,958 ± 65 (AA-100294)	Charcoal	1	1		Х	Х
P: 11,020 ± 140 (Beta-75225)	charcoal	1	1	Х		Х

Table S7.2. YDB_{Obs} measurements retained in each subsample (marked with 'X's). Alphabetical site IDs correspond to Fig. 1 and Table 1 in the main text.



Figure S7.1. Results of simulations with YDB_{Obs} subsampled to 19 measurements. Blue geometry corresponds to LST simulations, and red geometry corresponds to YDB simulations. Refer to Fig. 4 (main text) for interpreting the segments and bands in each panel.

8. References for Supplemental Information Appendix

- Aura Tortosa, J.E., Miret i Estruch, C., Morales Pérez, J.V., 2008. Coves de Santa Maira (Castell de Castells, La Marina Alta, Alacant). Campaña de 2008. *Saguntum, PLAV* 40, 227–232.
- Baales, M., Bittmann, F., Kromer, B., 1998. Verkohlte Bäume im Trass der Laacher See-Tephra bei Kruft (Neuwieder Becken): Ein Beitrag zur Datierung des Laacher See-Ereignisses und zur Vegetation der Allerød-Zeit am Mittelrhein. Archäologisches Korrespondenzblatt 28, 191–204.
- Baales, M., Jöris, O., Street, M., Bittmann, F., Weninger, B., Wiethold, J., 2002. Impact of the Late Glacial Eruption of the Laacher See Volcano, Central Rhineland, Germany. *Quaternary Research* 58, 273–288.
- Bement, L.C., Madden, A.S., Carter, B.J., Simms, A.R., Swindle, A.L., Alexander, H.M., Fine, S., Benamara, M., 2014. Quantifying the distribution of nanodiamonds in pre-Younger Dryas to recent age deposits along Bull Creek, Oklahoma Panhandle, USA. *Proceedings of the National Academy of Sciences USA* 111, 1726–1731. https://doi.org/10.1073/pnas.1309734111
- Bengtsson, H., Bravo, H.C., Gentleman, R., Hossjer, O., Jaffee, H., Jian, D., Langfelder, P., Hickey, P., 2018. matrixStats: Functions that Apply to Rows and Columns of Matrices (and to Vectors).
- Bergin, K.A., 2011. The Archaeology of the Talega Site (CA-ORA-907), Orange County, California: Perspective on the Prehistory of Southern California. Viejo California Associates, Mission Viejo. Prepared for the District of the US Army Corps of Engineers, Los Angeles.
- Bevan, A., Crema, E.R., 2018. rcarbon v1.2.0: Methods for calibrating and analyzing radiocarbon dates.
- Boaretto, E., Bryant, C., Carmi, I., Cook, G., Gulliksen, S., Harkness, D., Heinemeier, J., McClure, J., McGee, E., Naysmith, P., 2003. How reliable are radiocarbon laboratories? A report on the Fourth International Radiocarbon Inter-comparison (FIRI)(1998–2001). *Antiquity* 77, 146–154.
- Bunch, T.E., Hermes, R.E., Moore, A.M.T., Kennett, D.J., Weaver, J.C., Wittke, J.H., DeCarli, P.S., Bischoff, J.L., Hillman, G.C., Howard, G.A., Kimbel, D.R., Kletetschka, G., Lipo, C.P., Sakai, S., Revay, Z., West, A., Firestone, R.B., Kennett, J.P., 2012. Very hightemperature impact melt products as evidence for cosmic airbursts and impacts 12,900 years ago. *Proceedings of the National Academy of Sciences USA* 109, E1903–E1912.
- Caudullo, G., de Rigo, D., 2016. *Populus tremula* in Europe: distribution, habitat, usage and threats, in: San-Miguel-Ayanz, J., de Rigo, D., Caudullo, G., Houston, T., Mauri, A. (Eds.), *European Atlas of Forest Tree Species*. Publication Office of the European Union, Luxembourg, pp. 138–139.
- Firestone, R.B., West, A., Kennett, J.P., Becker, L., Bunch, T.E., Revay, Z.S., Schultz, P.H., Belgya, T., Kennett, D.J., Erlandson, J.M., Dickenson, O.J., Goodyear, A.C., Harris, R.S., Howard, G.A., Kloosterman, J.B., Lechler, P., Mayewski, P.A., Montgomery, J., Poreda, R., Darrah, T., Hee, S.S.Q., Smith, A.R., Stich, A., Topping, W., Wittke, J.H., Wolbach, W.S., 2007. Evidence for an extraterrestrial impact 12,900 years ago that contributed to the megafaunal extinctions and the Younger Dryas cooling. *Proceedings of the National Academy of Sciences USA* 104, 16016–16021.
- Frechen, J., 1952. Die Herkunft der spätglazialen Bimstuffe in mittel-und süddeutschen Mooren. *Geologisches Jahrbuch* 67, 209–230.
- Frechen, J., 1959. Die Tuffe des Laacher Vulkangebietes als quartärgeologische Leitgesteine und Zeitmarken. *Fortschritte der Geologie in Rheinland und Westfalen* 4, 363–70.

- Goodyear, A.C., 2013. Update on the 2012–2013 activities of the Southeastern Paleoamerican Survey. *Legacy* 17, 10–12.
- Hajdas, I., Ivy, S.D., Beer, J., Bonani, G., Imboden, D., Lotted, A.F., Sturm, M., Suter, M., 1993. AMS radiocarbon dating and varve chronology of Lake Soppensee: 6000 to 12000 14C years BP. *Climate Dynamics* 9, 107–116.
- Hajdas, I., Ivy-Ochs, S.D., Bonani, G., Loiter, A.F., Zolitschka, B., Schlüchter, C., 1995. Radiocarbon age of the Laacher See tephra: 11,230 ± 40 BP. *Radiocarbon* 37, 149–154.
- Hajic, E.R., Mandel, R.D., Ray, J.H., Lopinot, N.H., 2007. Geoarchaeology of stratified Paleoindian deposits at the Big Eddy site, Southwest Missouri, U.S.A. *Geoarchaeology* 22, 891–934.
- Haynes, C.V., 2007. Radiocarbon dating at Murray Springs and Curry Draw, in: Haynes, C.V., Huckell, B.B. (Eds.), *Murray Springs: A Clovis Site with Multiple Activity Areas in the San Pedro Valley, Arizona*. University of Arizona Press, pp. 229–239.
- Haynes, V., Agogino, G., 1960. *Geological Significance of a New Radiocarbon Date from the Lindenmeier Site*. Denver Museum of Natural History Proceedings No. 9. Denver Museum of Natural History.
- Hedges, R.E.M., Housley, R.A., Ramsey, C.B., Klinken, G.J.V., 1993. Radiocarbon dates from the Oxford AMS system: Archaeometry Datelist 16. *Archaeometry* 35, 147–167.
- Heine, K., 1993. Warmzeitliche Bodenbildung im Bölling/Alleröd im Mittelrheingebiet. *Decheniana* 146, 315–324.
- International Study Group, 1982. An inter-laboratory comparison of radiocarbon measurements in tree rings. *Nature* 298, 619–623.
- Johnson, E., Holliday, V.T., 1997. Analysis of Paleoindian bonebeds at the Clovis Site: New data from old excavations. *Plains Anthropologist* 42, 329–352.
- Kennett, D., Kennett, J., West, G., Erlandson, J., Johnson, J., Hendy, I., West, A., Culleton, B., Jones, T., Staffordjr, T., 2008. Wildfire and abrupt ecosystem disruption on California's Northern Channel Islands at the Ållerød–Younger Dryas boundary (13.0–12.9ka). *Quaternary Science Reviews* 27, 2530–2545.
- Kennett, J.P., Kennett, D.J., Culleton, B.J., Aura Tortosa, J.E., Bischoff, J.L., Bunch, T.E., Daniel, I.R., Erlandson, J.M., Ferraro, D., Firestone, R.B., Goodyear, A.C., Israde-Alcántara, I., Johnson, J.R., Jordá Pardo, J.F., Kimbel, D.R., LeCompte, M.A., Lopinot, N.H., Mahaney, W.C., Moore, A.M.T., Moore, C.R., Ray, J.H., Stafford, T.W., Tankersley, K.B., Wittke, J.H., Wolbach, W.S., West, A., 2015. Bayesian chronological analyses consistent with synchronous age of 12,835–12,735 Cal B.P. for Younger Dryas boundary on four continents. *Proceedings of the National Academy of Sciences USA* 112, E4344–E4353.
- Kinzie, C.R., Que Hee, S.S., Stich, A., Tague, K.A., Mercer, C., Razink, J.J., Kennett, D.J., DeCarli, P.S., Bunch, T.E., Wittke, J.H., Israde-Alcántara, I., Bischoff, J.L., Goodyear, A.C., Tankersley, K.B., Kimbel, D.R., Culleton, B.J., Erlandson, J.M., Stafford, T.W., Kloosterman, J.B., Moore, A.M.T., Firestone, R.B., Aura Tortosa, J.E., Jordá Pardo, J.F., West, A., Kennett, J.P., Wolbach, W.S., 2014. Nanodiamond-Rich Layer across Three Continents Consistent with Major Cosmic Impact at 12,800 Cal BP. *Journal of Geology* 122, 475–506.
- Kromer, B., Spurk, M., Remmele, S., Barbetti, M., Joniello, V., 1998. Segments of Atmospheric 14C Change as Derived from Late Glacial and Early Holocene Floating Tree-Ring Series. *Radiocarbon* 40, 351–358. https://doi.org/10.1017/S0033822200018221
- Lopinot, N.H., Ray, J.H., Conner, M.D., 1998. The 1997 Excavations at the Big Eddy Site (23CE426) in Southwest Missouri. Southwest Missouri State University, Springfield, Missouri.
- Lopinot, N.H., Ray, J.H., Connor, M.D., 2000. *The 1999 Excavations at the Big Eddy Site* (23CE426). Center for Archaeological Research Special Publication. Southwest Missouri State University, Springfield, Missouri.

McDonald, L., Chivall, D., Miles, D., Bronk Ramsey, C., 2019. Seasonal variations in the ¹⁴C content of tree rings: Influences on radiocarbon calibration and single-year curve construction. *Radiocarbon* 61, 185–194.

McElreath, R., 2017. rethinking v1.59: Statistical Rethinking book package.

Moore, A.M.T., Hillman, G.C., Legge, A.J., Huxtable, J., 2000. Village on the Euphrates: from Foraging to Farming at Abu Hureyra. Oxford University Press, Oxford.

- Pedersen, T.L., 2018. patchwork: The Composer of ggplots.
- R Core Team, 2018. R: A language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria.
- Reimer, P.J., Bard, E., Bayliss, A., Beck, J.W., Blackwell, P.G., Ramsey, C.B., Buck, C.E., Cheng, H., Edwards, R.L., Friedrich, M., 2013. IntCal13 and Marine13 radiocarbon age calibration curves 0–50,000 years cal BP. *Radiocarbon* 55, 1869–1887.
- Rubin, M., Alexander, C., 1960. U. S. Geological Survey radiocarbon dates V. *Radiocarbon* 2, 129–185.
- Schweitzer, H.-J., 1958. Entstehung und Flora des Trasses im nördlichen Laachersee-Gebiet. *E&G Quaternary Science Journal* 9, 28–48.
- Scott, E.M., Aitchison, T.C., Harkness, D.D., Cook, G.T., Baxter, M.S., 1990. An Overview of All Three Stages of the International Radiocarbon Intercomparison. *Radiocarbon* 32, 309– 319.
- Scott, E.M., Cook, G.T., Naysmith, P., 2010a. The Fifth International Radiocarbon Intercomparison (VIRI): An assessment of laboratory performance in Stage 3. *Radiocarbon* 52, 859–865.
- Scott, E.M., Cook, G.T., Naysmith, P., 2010b. A report on Phase 2 of the Fifth International Radiocarbon Intercomparison (VIRI). *Radiocarbon* 52, 846–858.
- Scott, E.M., Cook, G.T., Naysmith, P., 2007. Error and uncertainty in radiocarbon measurements. Radiocarbon 49, 427–440.
- Scott, E.M., Cook, G.T., Naysmith, P., Bryant, C., O'Donnell, D., 2007. A report on Phase 1 of the 5th International Radiocarbon Intercomparison (VIRI). *Radiocarbon* 49, 409–426.
- Scott, E.M., Harkness, D.D., Cook, Gt., 1998. Interlaboratory comparisons: lessons learned. *Radiocarbon* 40, 331–340.
- Stan Development Team, 2018. RStan 2.17.2: The R interface for Stan.
- Street, M.J., 1993. Analysis of Late Palaeolithic and Mesolithic Faunal Assemblages in the Northern Rhineland, Germany (Ph.D. dissertation). University of Birmingham.
- Tankersley, K.B., Redmond, B.G., 1999. Radiocarbon dating of a Paleoindian projectile point from Sheriden Cave, Ohio. *Current Research in the Pleistocene* 16, 76–77.
- van den Bogaard, P., Schmincke, H.-U., 1985. Laacher See Tephra: A widespread isochronous late Quaternary tephra layer in central and northern Europe. *Geological Society of America Bulletin* 96, 1554.
- van Geel, P., Coope, G.R., van der Hammen, T., 1989. Palaeoecology and stratigraphy of the late glacial type section at Usselo (The Netherlands). *Review of Paleobotany and Palynology* 60, 25–129.
- van Hoesel, A., Hoek, W.Z., Braadbaart, F., van der Plicht, J., Pennock, G.M., Drury, M.R., 2012. Nanodiamonds and wildfire evidence in the Usselo horizon postdate the Allerød-Younger Dryas boundary. *Proceedings of the National Academy of Sciences USA* 109, 7648–7653.
- von Wühlisch, G., 2009. *EUFORGEN Technical Guidelines for genetic conservation and use for Eurasian aspen (Populus tremula)*. Biodiversity International, Rome.
- Walton, A., Trautman, M.A., Friend, J.P., 1961. Isotopes, Inc. Radiocarbon Measurements I. *Radiocarbon* 3, 47–59.
- Waters, M.R., Stafford, T.W., Redmond, B.G., Tankersley, K.B., 2009. The age of the Paleoindian assemblage at Sheriden Cave, Ohio. *American Antiquity* 74, 107–111.

Wickham, H., 2016. ggplot2: Elegant Graphics for Data Analysis. Springer-Verlag, New York. Wickham, H., 2007. Reshaping data with the reshape package. *Journal of Statistical Software* 21, 1–20.

Wittke, J.H., Weaver, J.C., Bunch, T.E., Kennett, J.P., Kennett, D.J., Moore, A.M.T., Hillman, G.C., Tankersley, K.B., Goodyear, A.C., Moore, C.R., Daniel, I.R., Ray, J.H., Lopinot, N.H., Ferraro, D., Israde-Alcántara, I., Bischoff, J.L., DeCarli, P.S., Hermes, R.E., Kloosterman, J.B., Revay, Z., Howard, G.A., Kimbel, D.R., Kletetschka, G., Nabelek, L., Lipo, C.P., Sakai, S., West, A., Firestone, R.B., 2013. Evidence for deposition of 10 million tonnes of impact spherules across four continents 12,800 y ago. *Proceedings of the National Academy of Sciences USA*110, E2088–E2097.